

I Earth's Interior

- A How we know what's inside
 - i) Astronomy - Physics
 - (1) Calculating Earth's Density
 - (2) Gravitational influence, mass
 - (3) Volume, Shape, Diameter
 - (4) calc. Density of Earth 5.5 g/cm^3
 - (5) Surface 2.8 g/cm^3
 - ii) Drilling through the crust – deepest is 12 km (≈ 7.4 miles)
 - iii) Earth's delivery service
 - (1) Kimberlite Pipes and Xenoliths
 - (2) Volcanoes
 - (3) Seismic data – (will discuss later)
 - iv) Vibration energy waves
 - v) Earthquake data
- B Interior Structure
 - i) Chemical Composition
 - (1) Crust
 - (a) 5 to 75 km thick
 - (b) Solid
 - (c) Silicon (Si) and Oxygen (O)
 - (d) Continental
 - (i) 5 to 8 km thick
 - (ii) Old (billions of years in age)
 - (iii) Deformed
 - (iv) Density 2.7 g/cm^3 - Granite
 - (e) Oceanic
 - (i) 5 to 8 km thick
 - (ii) Relatively young (less than 200 million years)
 - (iii) Undeformed
 - (iv) Density 3.0 g/cm^3 - Basalt
 - (2) Mantle
 - (a) Approximately 2900 km thick
 - (b) 82% of Earth's Volume
 - (c) 68% of Earth's Mass
 - (d) Density increases with depth through the mantle
 - (e) 3.2 g/cm^3 in upper portion
 - (f) 5 g/cm^3 in lower portion
 - (g) Silicon (Si) and Oxygen (O) with some iron (Fe) and magnesium (Mg)
 - (3) Core
 - (a) Total diameter – 7000 km
 - (b) 16% of Earth's Volume
 - (c) 32% of Earth's Mass – Density 10.8 g/cm^3
 - (d) Iron (Fe) Rich
 - ii) Earth's delivery service
- C Interior Components - Physical Properties
 - i) Lithosphere

- (1) Solid, strong, and rigid outer layer
- (2) Includes all of crust and the upper mantle
- (3) 10 to 300 km thick
- ii) Asthenosphere
 - (1) High Temperature
 - (2) Moderate Pressure
 - (3) Soft-plastic zone (Rock can flow)
 - (4) Upper Mantle
 - (5) 100 to 500 km thick
- iii) Mesosphere
 - (1) High Temperature Zone
 - (2) High Pressure Zone
 - (3) Rock is stronger and more rigid than in the overlying Asthenosphere
 - (4) Remainder of Mantle
- iv) A Two Component Core
 - (1) Inner Core
 - (a) \approx 1200 km thick
 - (b) Solid
 - (2) Outer Core
 - (a) \approx 2300 km thick
 - (b) Liquid – Molten
 - (c) May have Nickel (Ni) present

II Plate Tectonics

A Plates and Plate Boundaries

- 1 Plates
 - a Can be oceanic crust
 - b Can be continental crust
 - c Can be a combination of both
- 2 Divergent
- 3 Convergent
- 4 Transform

B Divergent Boundaries – Mid-Ocean Ridges (MOR)

- 1 Separates 2 tectonic plates
- 2 Tectonic plates move apart
- 3 New Oceanic Crust is generated
- 4 Average crustal generation – 5 cm/yr; varies between 2-20 cm/yr
- 5 Rifting
 - a Upwelling below continent
 - b Crust extends, thins, & weakens
 - c Extension causes faulting and volcanism
 - d Forms rift valley
 - e Spreading continues
 - f Rift becomes deeper and wider – creates a linear sea
 - g Further spreading creates an ocean basin
- 6 Ophiolite Complex
 - a Oceanic Crust

- b Marine Sediment
 - c Pillow Basalts
 - d Sheeted Dikes
 - e Gabbro
- 7 Earthquakes and Divergent Boundaries
- a Earthquakes are shallow, less than 10 km in depth
 - b Small magnitude (low energy)
 - c Narrow pattern along MOR
- C Convergent Boundaries
- 1 Plates come together
- a Ocean-Ocean Convergence
 - b Ocean-Continent Convergence
 - c Continent-Continent Convergence
- 2 Subduction Zones
- a Ocean-ocean collision or ocean-continent collision
 - b Density differences in crust forces one plate to descend under the other plate
 - c More dense plate descends at an angle between 15° to 75° with an average of $\approx 45^\circ$
 - d Subduction of plate creates a deep-ocean trench
 - i Range in depth between 8 – 12 km
 - ii 100's km wide and 1000's km long
- 3 Ocean-Continent Collision
- a Ocean crust is more dense (3.0 g/cm³) than continental crust (2.8 g/cm³)
 - b Ocean Crust subducts
 - c Within the mantle, plate begins to melt
 - d Melt has basaltic to intermediate composition
 - e Hot melt (magma) is less dense (more buoyant) than surrounding material and rises
 - f Ascending magma incorporates continental crust – composition becomes intermediate to felsic
 - g Creates linear zone of igneous activity (volcanoes or intrusions-plutons) called a Continental Arc
- 4 Ocean-Ocean Convergent Zone - Ocean-Ocean Collision
- a Similar to Ocean-Continent Convergent Zone except no influence of Continental Crust
 - b Magmas are basaltic to intermediate in composition
 - c Ascending magma creates linear Volcanic Island Arc
- 5 Continent-Continent Convergent Zone (Collision)
- a Is preceded by Ocean-Continent Convergence
 - b Density between plates is similar – neither will subduct
 - c Crust buckles, fractures, shortens, thickens due to compressional stress
 - d Creates large mountains of folded and faulted rocks.
- 6 Earthquakes & Convergent Boundaries
- a Earthquake depth ranges from shallow (< 10 km) to deep (> 300 km)

- b High magnitude (high energy) – 95% of all earthquake energy is released from convergent zone earthquakes
 - c Wide zone, but with a pattern
- D Transform Boundary
 - 1 Transform Faults
 - a Plates slide past one another
 - b Crust is neither produced nor destroyed
 - c Facilitate plate movement
 - d Connect segments of MOR
 - e Found in both ocean crust and continental crust
 - f Can create compressional structures (folds) and extensional structures (basins)
 - 2 Earthquakes and Transform Boundaries
 - a Earthquakes are shallow, less than 10 km in depth
 - b Small magnitude (low energy)
 - c Common
 - d Narrow pattern along MOR
- E Hot Spots
- F Plate Tectonics – Driving Mechanisms
 - 1 Driven by the release of internal energy in the form of heat
 - 2 Proposed mechanisms
 - a Convection cells
 - b Differences in density and forces generated by plates
 - c Forces Active on Plates
 - d Slab-pull – sinking of dense ocean crust pulls the plate
 - e Ridge-push – gravity pulls plate off of the MOR
 - f Basal drag – resistance of plate from underlying asthenosphere
 - g Mantel resistance – subducting plate is resisted in the asthenosphere and mesosphere
 - (a) Friction – resistance along transform faults and between converging plates